

# Basic Characteristics of Edge Turbulence with Zonal Flow and Global Alfvén Interactions

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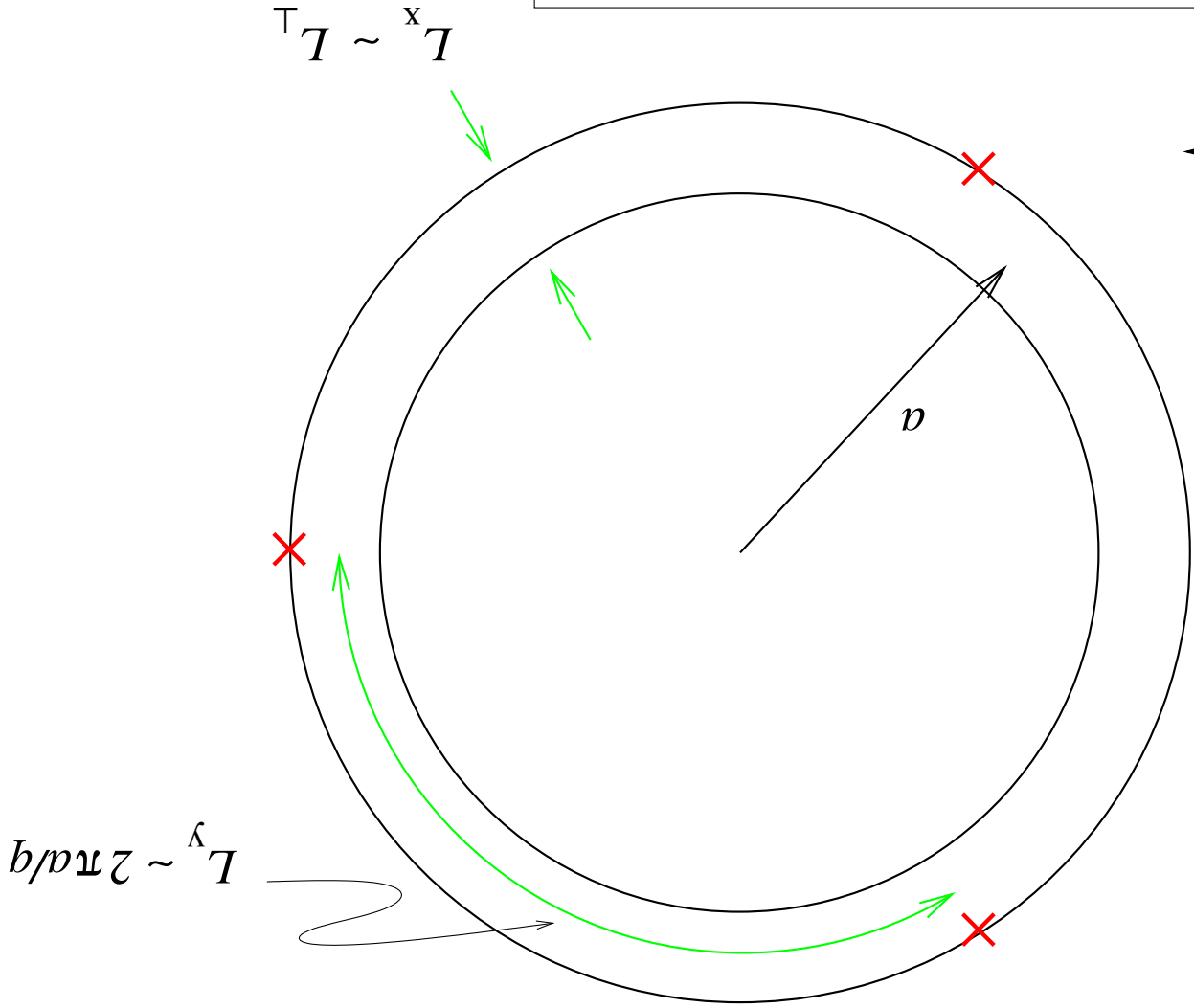
*Feb 2004*



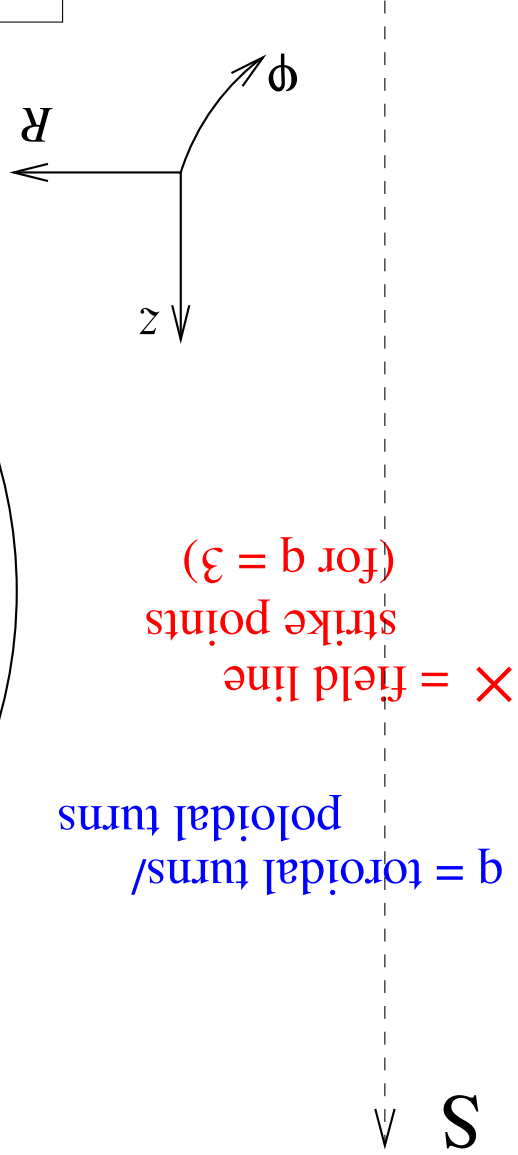
# Outline

- **Basic Ideas**
  - tokamak geometry, computational domain
  - basic dynamics, scales, computational resolution
  - gyrofluid model GEM3, treatment of various species
  - boundary conditions for “global” computation
- **Basic Results**
  - standard DALT turbulence
  - nonlinear self sustained drift wave instability
  - zonal flows and sidebands, global Alfvén transient  $\rightarrow$  P-S current (2D MHD equilibrium)
- **Specific Applications**
  - transport of a trace fluid (e.g., tritium)
  - SOL/edge coupling
- **Future Generalisation**
  - edge/core coupling, requires full inhomogeneity
  - hot trace ion adiabaticity (diamagnetic flows)

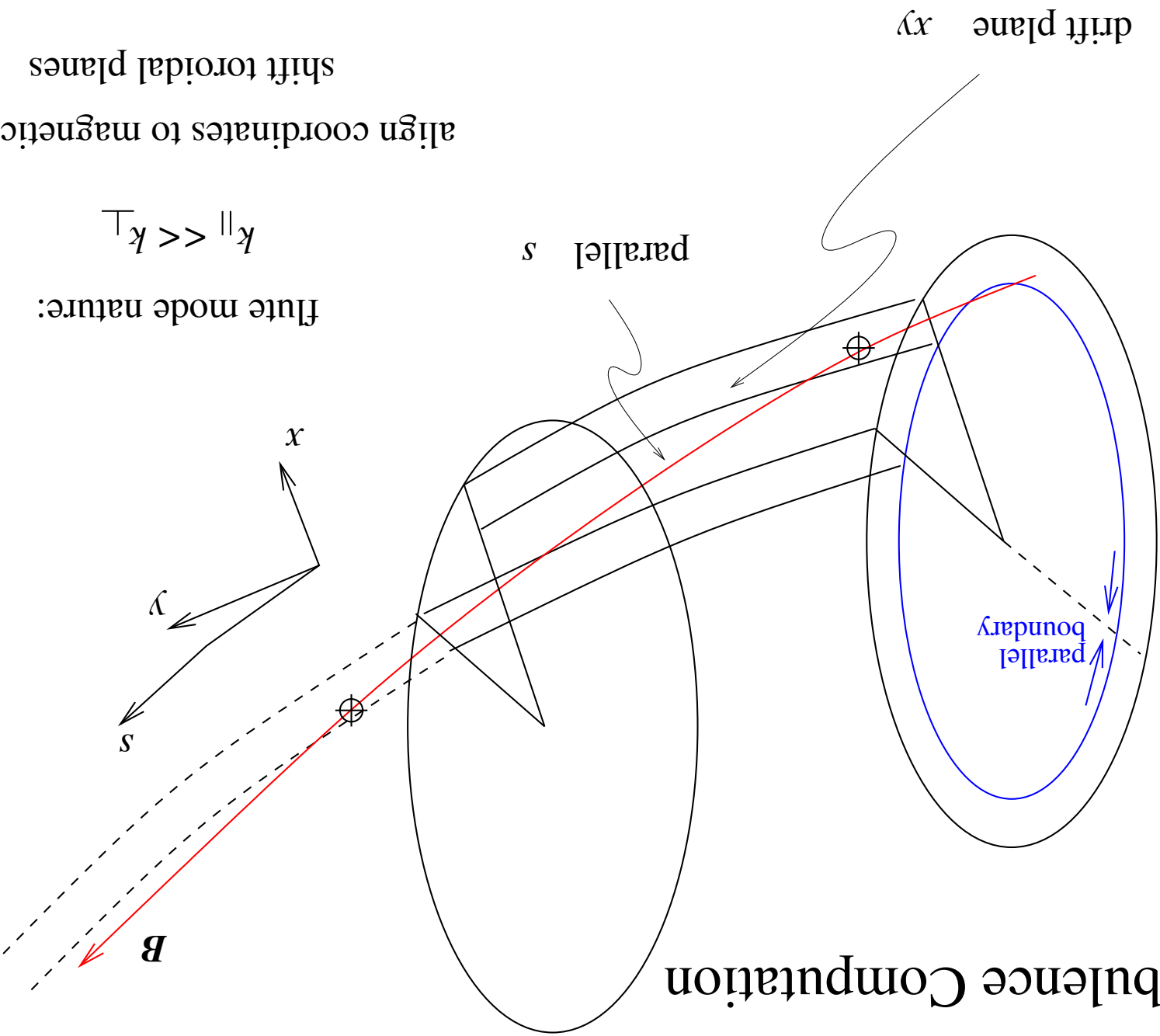
# Computational Domain



large aspect ratio:  $L^y \gg L^x$



# Turbulence Computation



# Magnetic Field

- toroidal coordinates  $(r, \theta, \zeta)$
- field aligned coordinate representation (simple version, normalised)
 
$$x = r - a \quad y_k = \frac{q_a}{a} [q(\theta) - s_k] - \zeta \quad s = \theta$$
- pitch parameter  $q(r) = B_\zeta / B_\theta$ , constants  $q_a = q(a)$  and  $s_k$
- globally consistent boundary conditions
  - preserves set of allowed  $k_\parallel$  values
  - required to capture field line connection (B Scott, Phys Plasmas 5 [1998] 2334)
- shifted metric: chosen so that  $g_{xy}^k = 0$  at  $s = s_k$ , hence ...
  - perp metric is unit diagonal at  $s = s_k$
  - $\partial/\partial s$  incurs shifts in  $y$  since  $y_{k\pm 1} \neq y_k$
  - required to capture slab modes (B Scott, Phys Plasmas 8 [2001] 447)
- normalised magnetic field strength is unity
- $\Delta_\perp^2$  and  $\mathbf{v}_E \cdot \nabla$  and  $\mathcal{K}$  involve only perp coordinates  $(x, y_k)$
- geometric quantities depend only upon parallel coordinate  $s$
- (all can be generalised to an arbitrary separatrix-free tokamak equilibrium)

## coordinates

aligned to mag field

reference q sweeps with  $q(r)$

require far fewer grid nodes

than for traditional methods

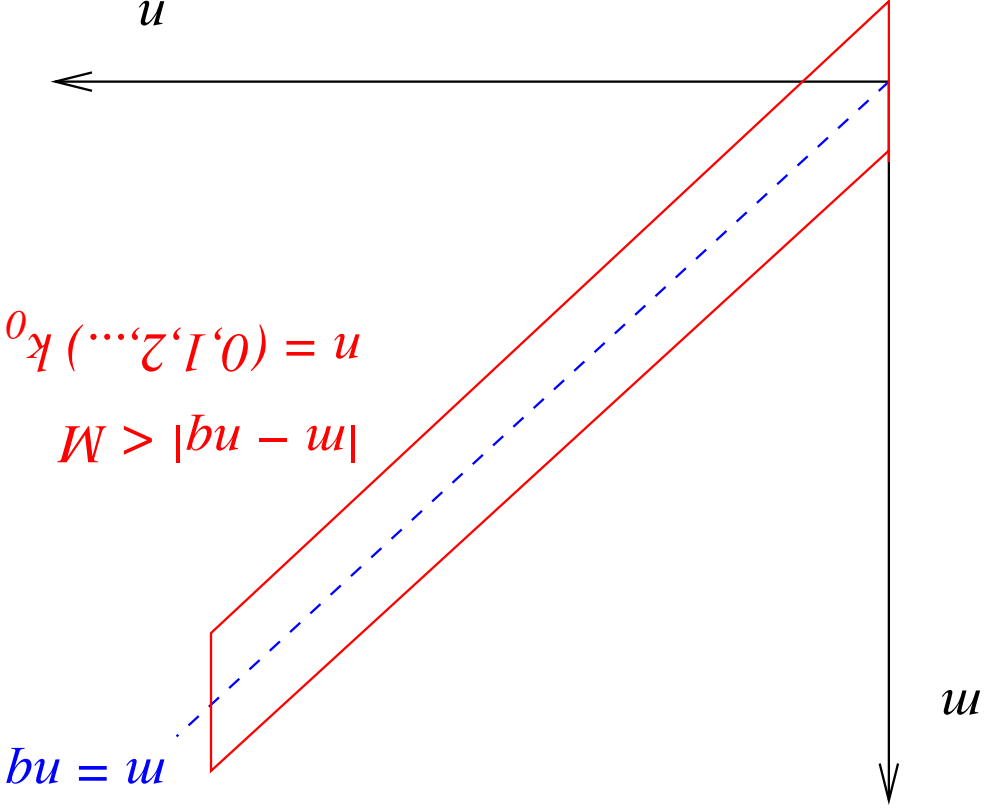
**scalability!**

resolve scales:

gradient to ion gyroradius

**require >100 toroidal modes!**

## mode space



global consistency

$$k_{\parallel} q R = m - n q$$

parallel grid nodes

$$M \gg n \text{ (delta } q)$$

# The GEM3 Model

- simplest electromagnetic gyrofluid model (B Scott, PFCF 45 [2003] A385)
- same ordering as in fluid drift model (often called “gyrokinetic”)
- two moments for each species: density, parallel velocity
- version incorporating background gradient into dependent variables
  - profile variation, self consistent transport equilibrium
  - however, still homogeneous equations

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- parameters and normalisation convention

- magnetic field representation and coordinates

- moment equations and polarisation

- full generalisation (not shown): six moments (incl  $T_{\perp}$  and  $T_{\parallel}$  and  $q_{\perp}$  and  $q_{\parallel}$ )

# Parameters

- gyro Bohm normalisation, to  $T_e$  and  $M_D$
- drift scale  $\rho_s$ , sound speed  $c_s$ , following “electron mobility, ion inertia”
- basic profile scale  $L_\perp$ , hence scale ratio “drift parameter”  $\delta = \rho_s/L_\perp \gg 1$
- field line connection length  $2\pi qR$  hence scale ratio  $qR/L_\perp \gg 1$
- low frequency drift turbulence scales

$$\rho_s = \frac{c}{eB} \sqrt{M_D T_e} \quad c_s = \sqrt{\frac{T_e}{M_D}} \quad \text{frequency } \frac{c_s}{L_\perp} \gg \Omega_i$$

- overall edge turbulence parameters (all  $> 1$  but not  $\gg 1$ )

$$\hat{\beta} = \frac{4\pi p_e}{B^2} \left( \frac{L_\perp}{qR} \right)^2 \quad \hat{\mu} = \frac{m_e}{M_D} \left( \frac{L_\perp}{qR} \right)^2 \quad C = \frac{0.51 v_e}{c_s} \frac{m_e}{M_D} \left( \frac{L_\perp}{qR} \right)^2$$

- background species parameters

$$a_z = \frac{Z n_e}{Z n_z} \quad \tau_z = \frac{Z T_e}{T_z} \quad \rho_z^2 = \frac{M_z T_z}{Z^2 M_D T_e} \quad \epsilon_z = \frac{M_z}{Z M_D} \left( \frac{L_\perp}{qR} \right)^2$$



# Moment Equations for Each Species, Polarisation

- density moment with source

$$\frac{dn_z}{dt} = -B \Delta_{\parallel} n_z + \kappa \left( \tilde{\phi}_G + \tau n_z \right) + S$$

- velocity moment

$$\beta \frac{\partial A_{\parallel}}{\partial t} + \epsilon_z \frac{dn_z}{dt} + C J_{\parallel} = -\Delta_{\parallel} \left( \tilde{\phi}_G + \tau n_z \right) + \epsilon_z \tau \kappa (2n_z)$$

- gyroaveraged potential, species specific

$$\tilde{\phi}_G = \Gamma_{1/2}^0(\tilde{\phi}), \quad \text{with} \quad \Gamma_0 = \Gamma_0(k_{\perp}^2 \rho_z^2)$$

- polarisation: Poisson (quasineutrality) and Ampere  
 o species moments  $\rightarrow$  potentials, only place species communicate

$$\sum_z a_z \left[ \Gamma_{1/2}^0 n_z + \frac{\tau_z}{\Gamma_0 - 1} \tilde{\phi} \right] = 0 \quad - \quad \sum_z a_z \tilde{n}_z = \Delta_{\perp}^2 A_{\parallel}$$

# Boundary Conditions

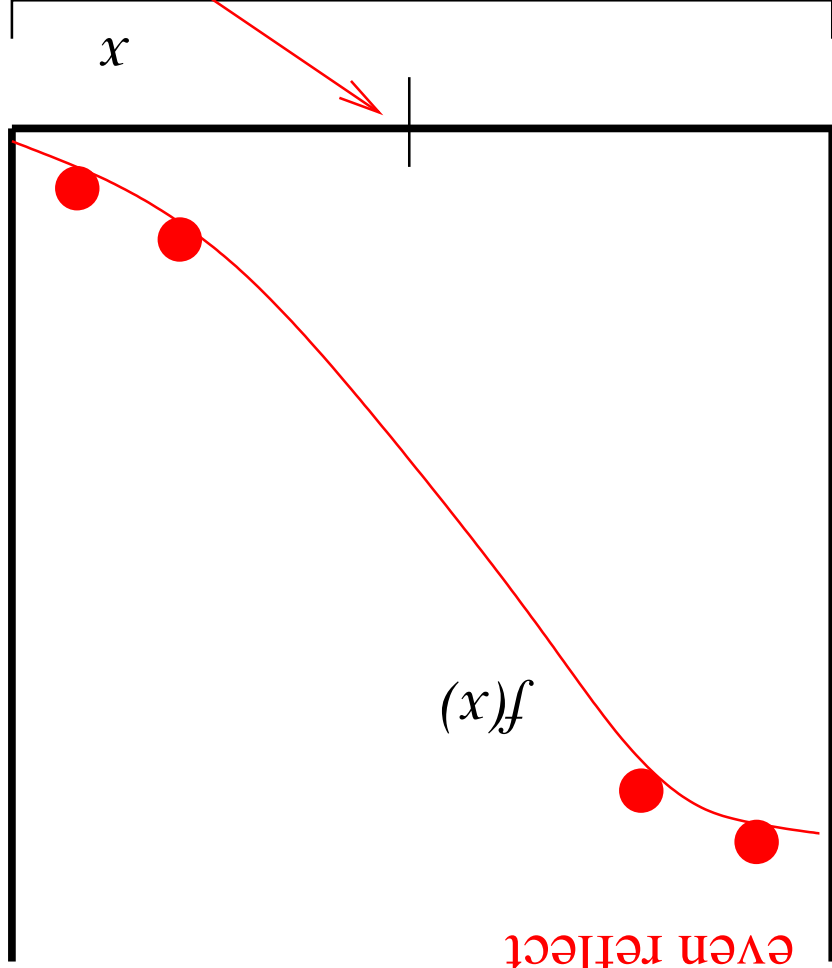
$x$  Neumann/Dirichlet

$y$  periodic

$s$  pseudo periodic (shifted metric)

odd reflect

even reflect



$x$

$L_x$

2 guard cells req'd for derivs

define parms at this location

# GEM3 Energetics

- very similar to fluid version, things in different places
  - role of polarisation: ExB energy looks like a charge potential energy
- potential (ExB) energy ( $T_0^{1/2}$  is Hermitian ...)

$$\frac{\partial}{\partial t} \sum^z a_z \frac{\partial \phi_z}{\partial n_z} = \frac{1 - \Gamma_0^z}{2} \sum^z a_z \frac{\partial \phi_z}{\partial t}$$

- thermal free energy (local eqs: like entropy)

$$\frac{\partial}{\partial t} \sum^z a_z T_z \frac{\partial n_z}{\partial n_z} = \frac{1}{2} \sum^z a_z T_z \frac{\partial n_z}{\partial t}$$

- parallel kinetic energy

$$\frac{\partial}{\partial t} \sum^z a_z \frac{\partial n_z}{\partial n_z} = \frac{1}{2} \sum^z a_z \frac{\partial n_z}{\partial t}$$

- magnetic energy

$$\frac{\partial}{\partial t} \sum^z a_z \frac{\partial \tilde{A}_{\parallel}}{\partial \tilde{A}_{\parallel}} = \frac{1}{2} \sum^z a_z \frac{\partial \tilde{A}_{\parallel}}{\partial t}$$

# Conservation

- FxB and thermal free energy combine ...

$$\sum^z a^z \left( \tilde{\phi}^z + \tau^z \tilde{u}^z \right) \frac{\partial \tilde{u}^z}{\partial t} = \sum^z a^z \left( \tilde{\phi}^z + \tau^z \tilde{u}^z \right) \left[ -B \Delta_{\parallel} \frac{\tilde{u}^z}{B} + \kappa \left( \tilde{\phi}^z + \tau^z \tilde{u}^z \right) + S \right]$$

- magnetic and parallel kinetic energy combine ...

$$\sum^z a^z \tilde{u}^z_{\parallel} \left( \tilde{\beta} \frac{\partial \tilde{A}_{\parallel}}{\partial t} + \epsilon_s \frac{\partial \tilde{u}^z_{\parallel}}{\partial t} \right) = \sum^z a^z \tilde{u}^z_{\parallel} \left[ -\Delta_{\parallel} \left( \tilde{\phi}^z + \tau^z \tilde{u}^z \right) + \epsilon_z \tau^z \kappa \left( 2 \tilde{u}^z_{\parallel} \right) \right] - C \tilde{J}_{\parallel}^2$$

- transfer and conservation is clear (force potential versus parallel flow)

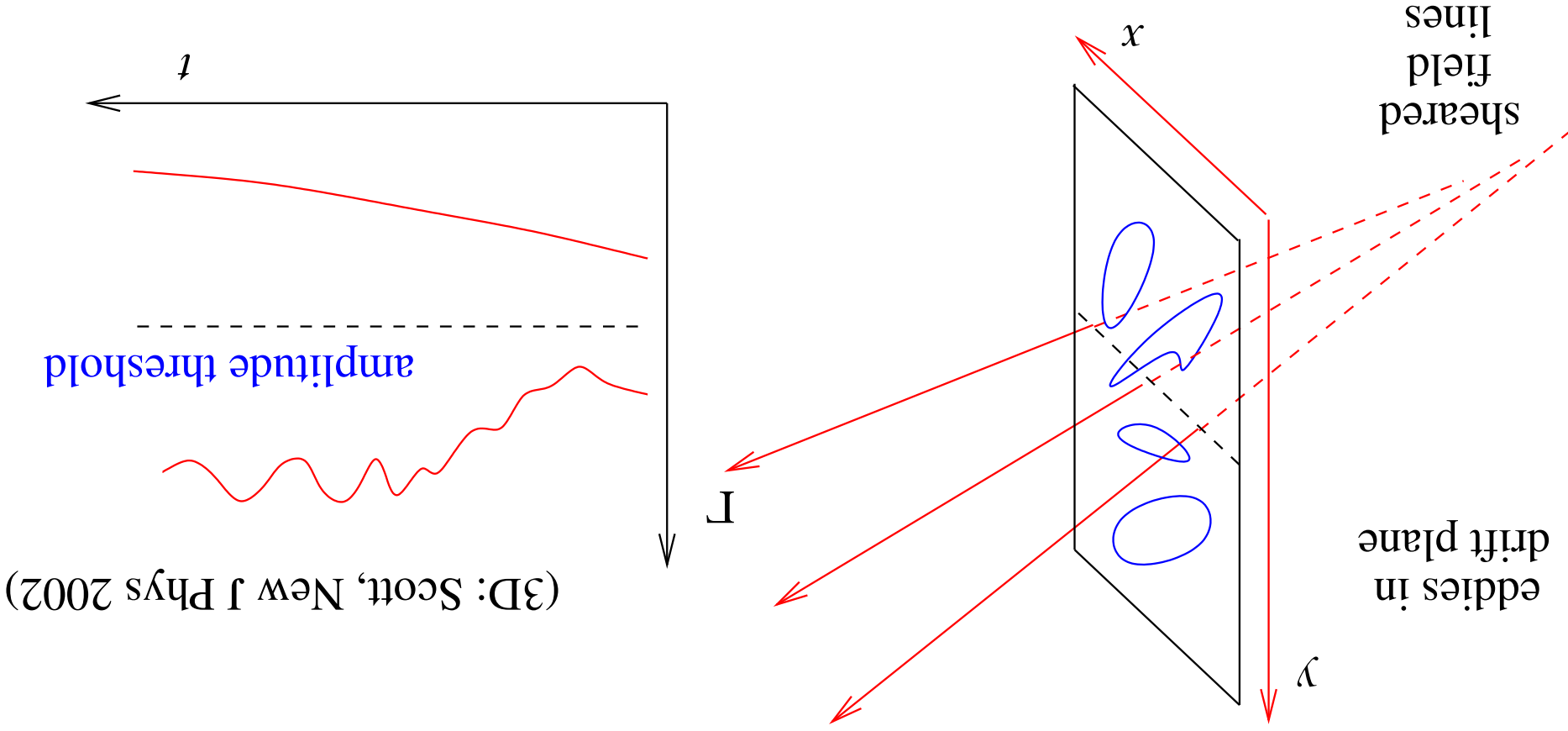
- zonal flow energy

$$\left\langle \frac{\partial \tilde{u}^z}{\partial t} \right\rangle \left\langle \tilde{\phi}^z \right\rangle \sum^z a^z = \frac{\tau^z}{1 - \Gamma_0} \frac{\partial \tilde{u}^z}{\partial t} \sum^z a^z \left\langle \tilde{\phi}^z \right\rangle^2$$

- further manipulations extract Reynolds, Maxwell stresses and geodesic coupling
  - vorticity is the same thing as “gyrofluid charge density”
  - etc, etc (not worked out here)

# Nonlinear Instability for Drift Waves

(Scott, Phys Rev Lett 1990; Phys Fl B 1992)



(3D: Scott, New J Phys 2002)

nonlinear vorticity rips open adiabatic response

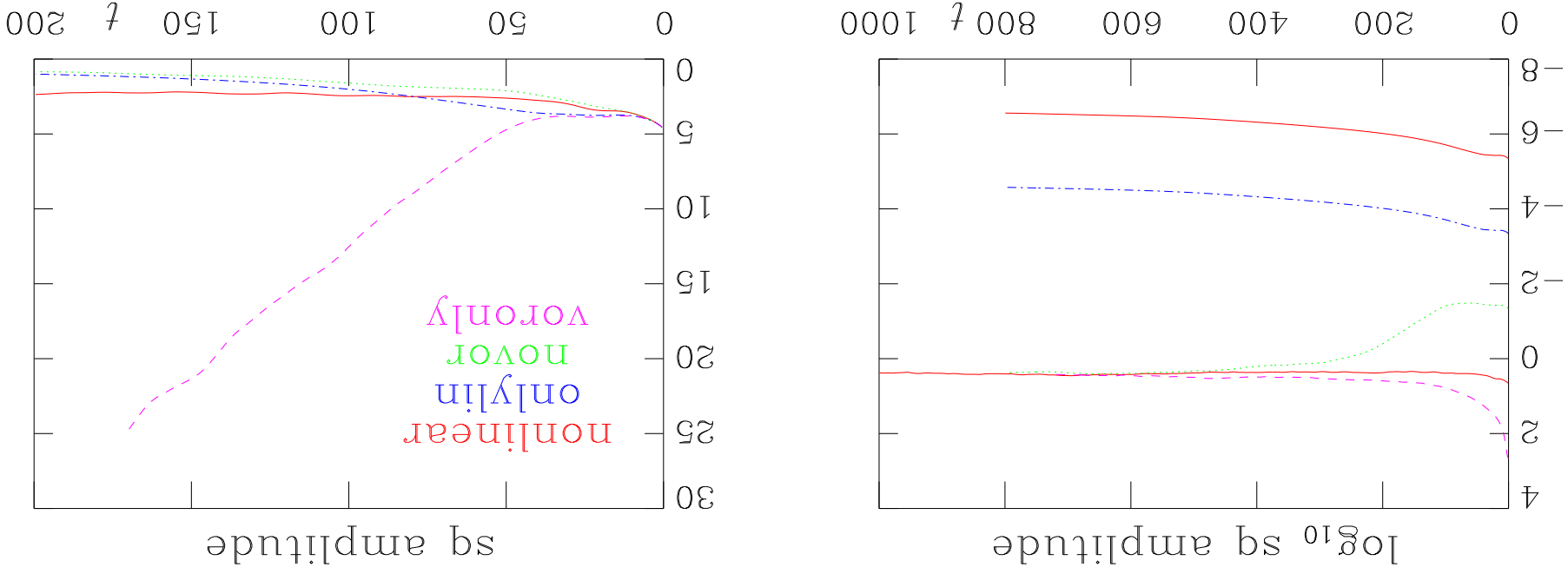
larger accessible space for parallel currents

larger phase shifts

—> nonlinear excitation/instability, ‘self sustained turbulence’

# self sustained turbulence

$C = 10, \beta = 0.1, 256 \times 256$

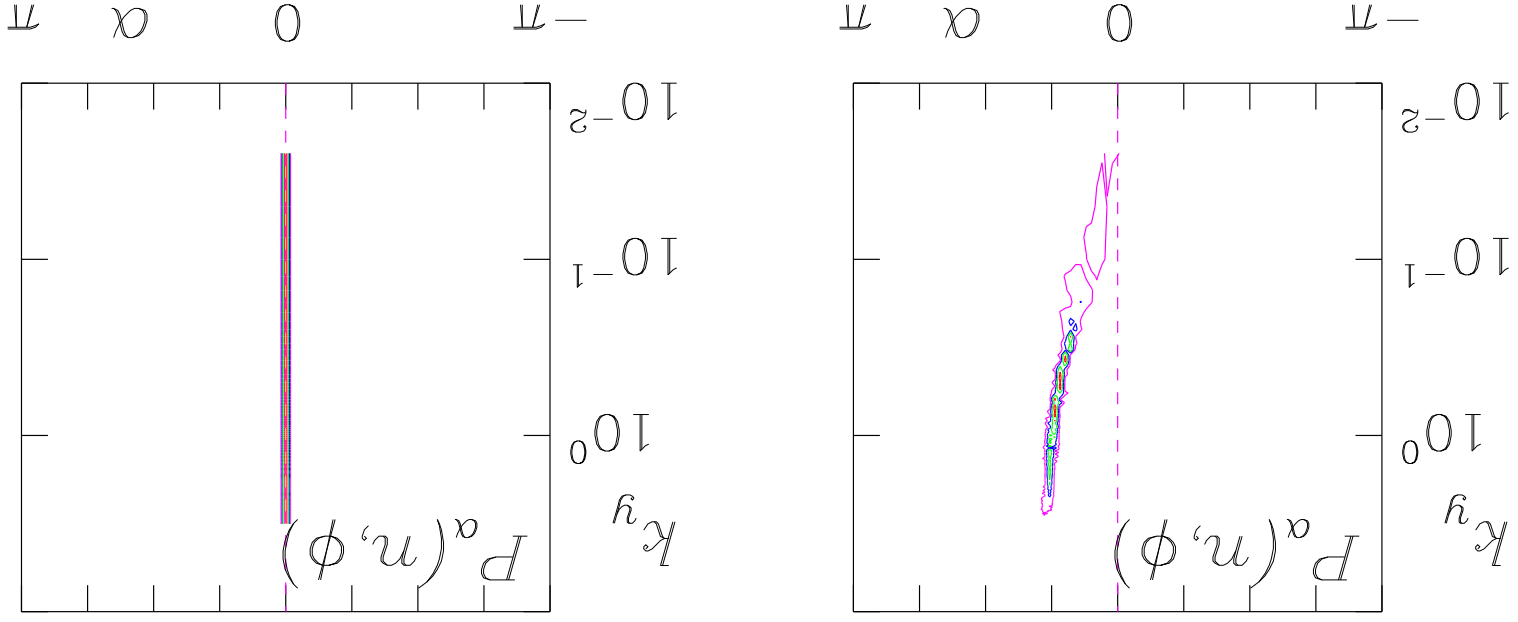


existence of amplitude threshold  
 convergence at late times (short correlation time)  
 $\mathbf{v}^E \cdot \Delta \tilde{U}$  is necessary and sufficient for self-sustainment  
 $\mathbf{v}^E \cdot \Delta \tilde{p}_e$  is necessary for saturation of nonlinear cases

# linear drift wave phase shifts

2D Hasegawa Wakatani,  $D = 0.01, 64 \times 256$

- linear start, all terms (left), no gradient (right)



- phase shift caused by gradient, part of dispersion relation
  - scalar function of perpendicular wavenumber

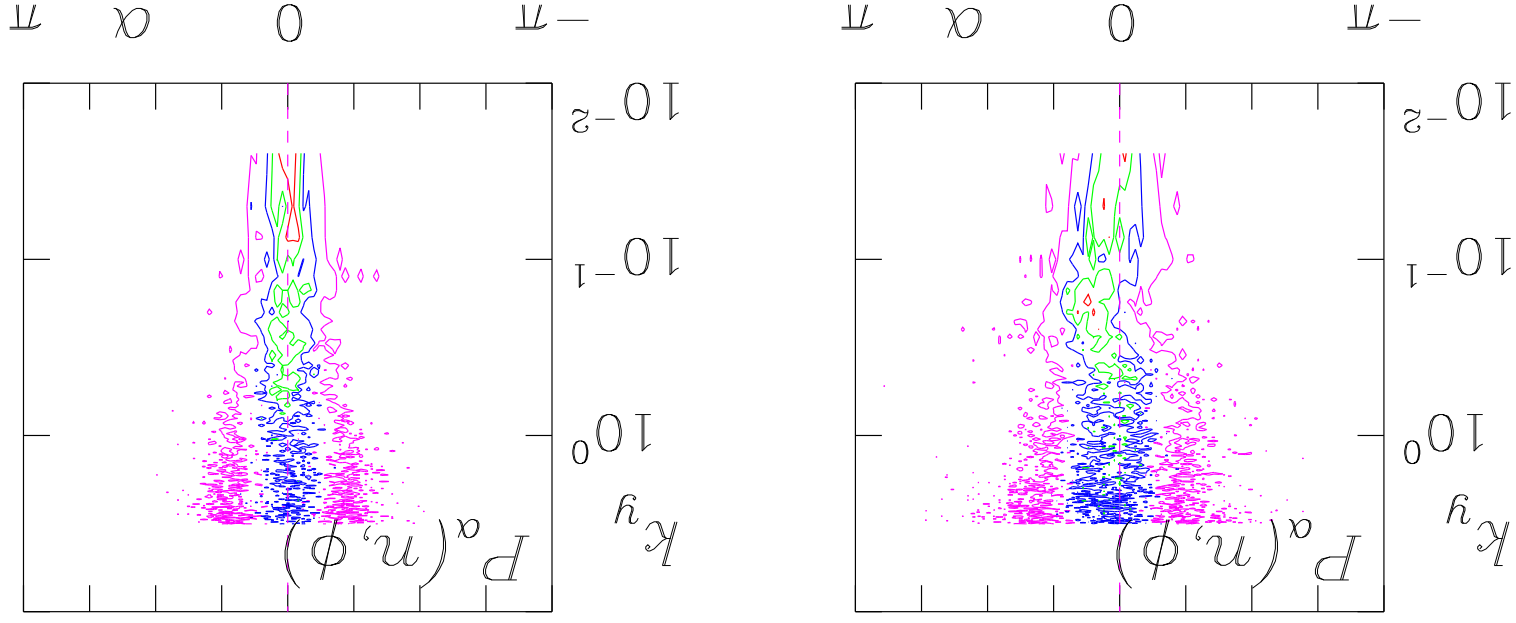
# drift wave turbulence phase shift distributions

2D Hasegawa Wakatani,  $D = 0.01, 64 \times 256$

- nonlinear start,

all terms (left),

no gradient (right)



- phase shifts caused by gradient, finds statistical equilibrium

- nonlinearities broaden distribution, gradient forcing selects for positive phase shifts

native drift wave vorticity overcomes linear growth rate  
mid range of spectrum does not sense action by linear instability



# toroidal case — linear start

compare linear and nonlinear mechanisms

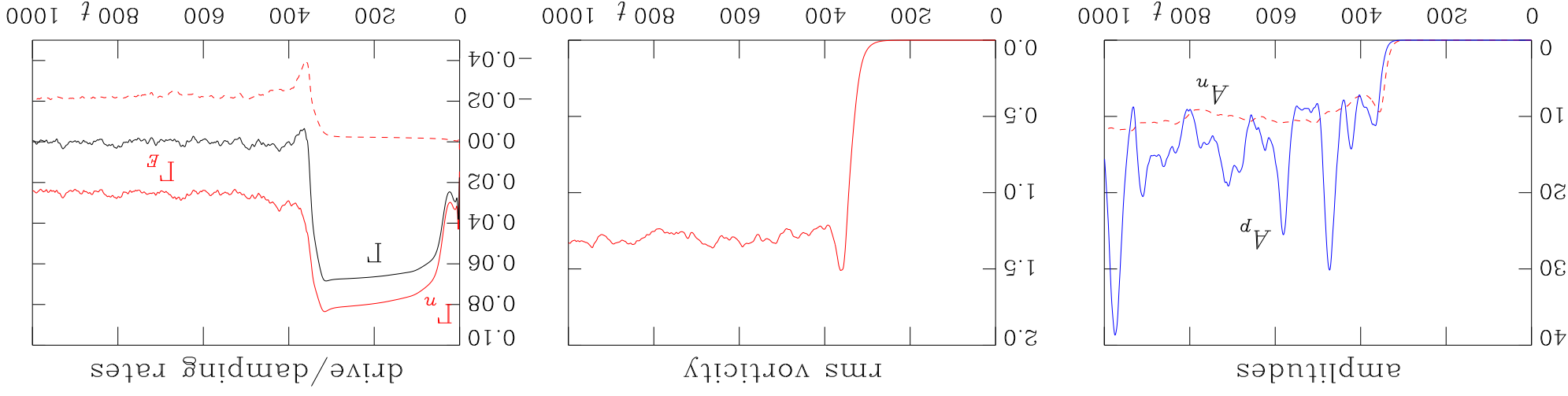
- start at small amplitude:  $\tilde{p}_e/p_e \sim 10^{-10}$

- linear eigenmode establishes itself — with clear ballooning mode signature

- transition to turbulence — saturation as  $\gamma \rightarrow 0$

- well developed turbulence at late times — statistical equilibrium

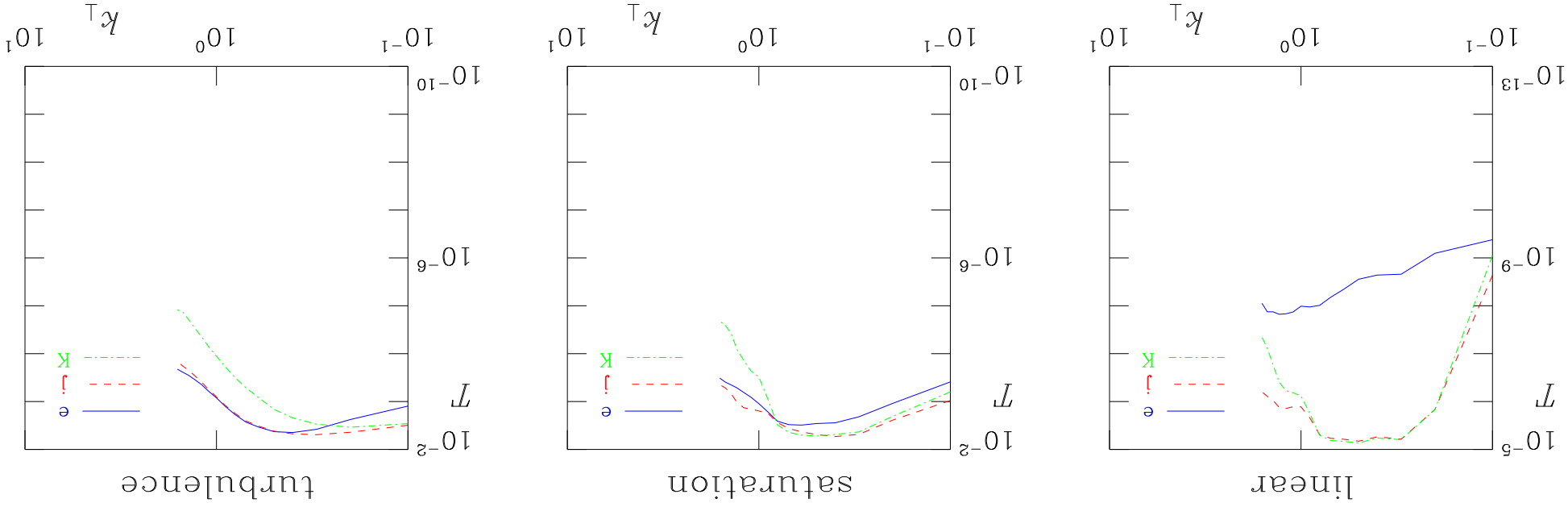
self sustained drift wave turbulence supersedes ballooning instability



# RMS transfer spectra — toroidal linear start

measures energy transfer in and out of ExB eddies at given scale

red: parallel currents, green: interchange forcing, blue: nonlinear vorticity scattering



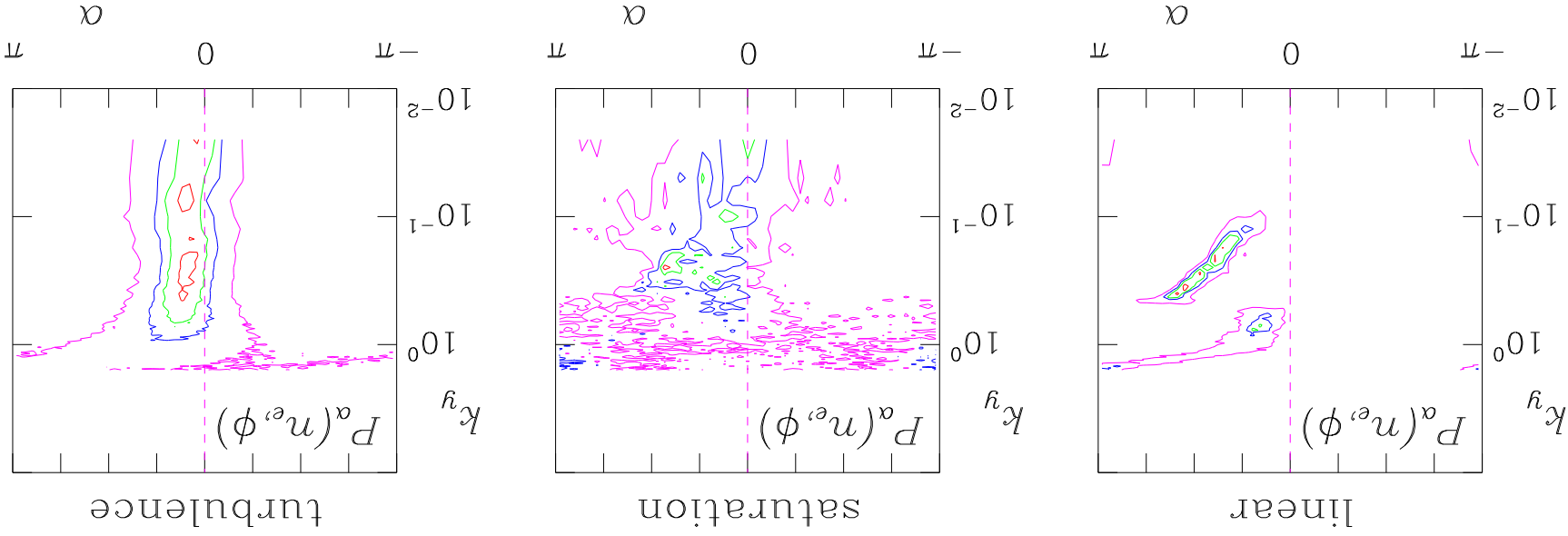
linear modes and fully developed turbulence have fundamentally different excitation mechanisms

# phase shift distributions — toroidal linear start

measures phase shift distributions for each wavelength

for linear modes: part of the dispersion relation

for turbulence: part of the statistical character involving damped or stable, as well as driven transients

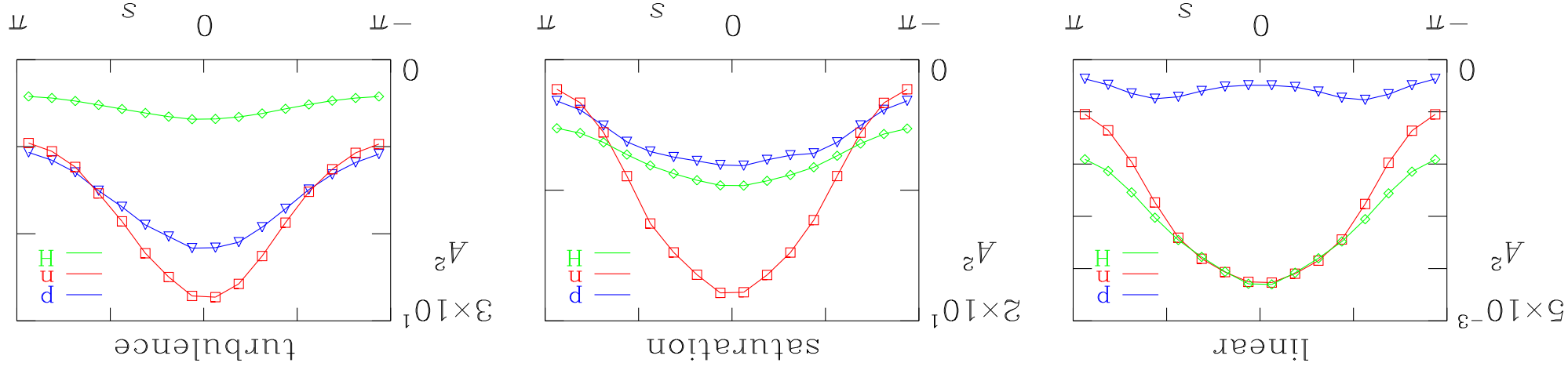


strong linear instability with large phase shift  
is no longer felt in the fully developed turbulence

# parallel mode structure — toroidal linear start

measures squared amplitudes as function of position along field lines

basic ballooning signal is ballooning in force potential,  $\tilde{h}_e = \tilde{n}_e - \tilde{\phi}$  (green line)



linear ballooning mode character is lost in fully developed turbulence

# Implications of the Nonlinear Instability

linear results cannot be *a priori* applied to edge turbulence

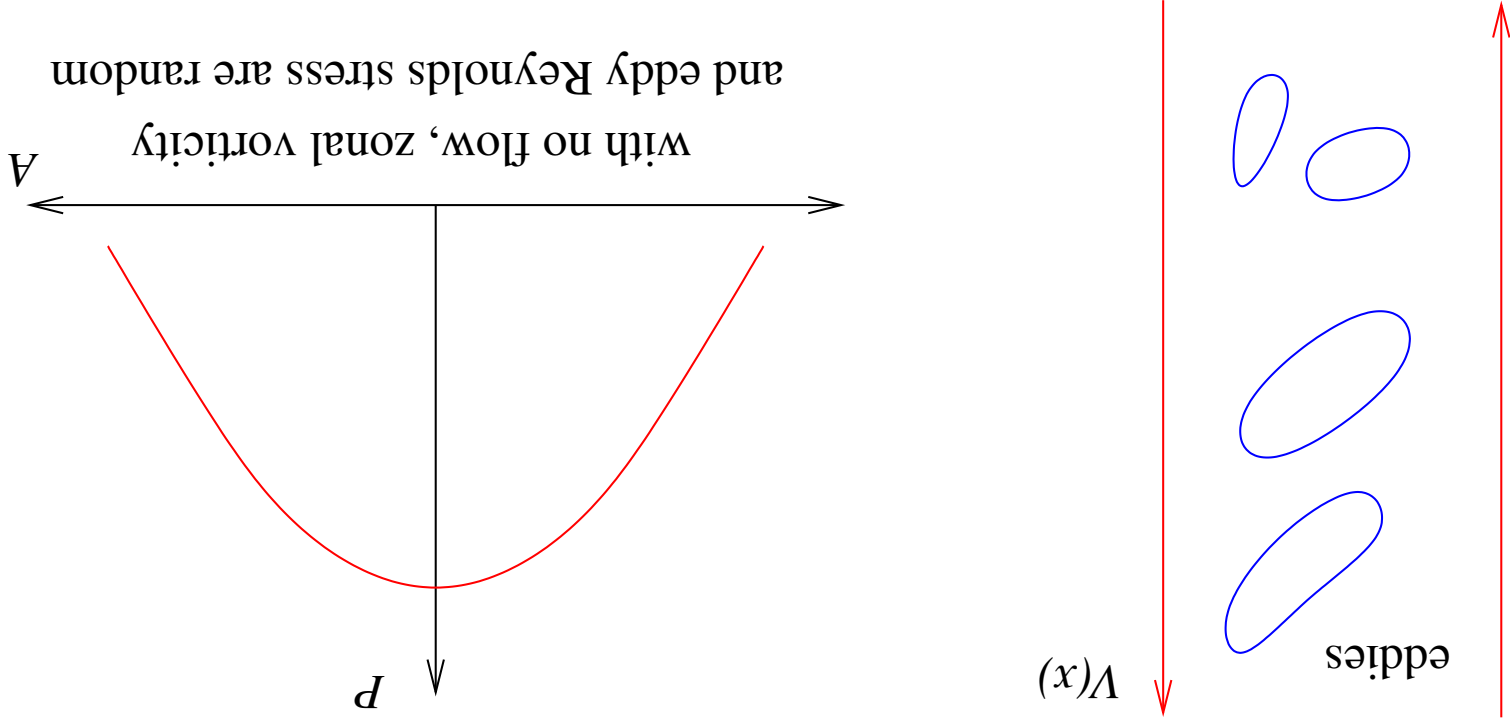
- no applicability for gamma/k scaling models
- linear mode relevance to be decided on case-by-case basis
- resistive ballooning wiped out by drift waves
- **ideal ballooning** is large scale  $\rightarrow$  still relevant (also peeling), but sets boundaries
- **most L-H models are ruled out** (linear  $\beta$  or  $C$  dependence)
- they *can* be tested by nonlinear code (homogeneous models)
- failing the test, the linear models are inapplicable to experiment

# Zonal ExB Flows

- zonal potential — flux surface averaged  $\tilde{\phi}$
- zonal flow is the ExB flow resulting from a zonal potential
- ExB flow is the advecting agent...
- a sheared ExB flow leads to reduced drive rates, transport
  - can suppress turbulence altogether if large enough
- shear suppression is conservative transfer from eddies to flow
  - Reynolds stress
- in toroidal geometry, zonal flow is compressible (quasistatically)
  - result of geodesic curvature
- “geodesic transfer effect” couples flows back to turbulence
  - (B Scott, Phys Letters A 320 [2003] 53)
- entire system reaches statistical equilibrium
  - flows are not “spun up” indefinitely by turbulence

# Reynolds Stress and Zonal Flows

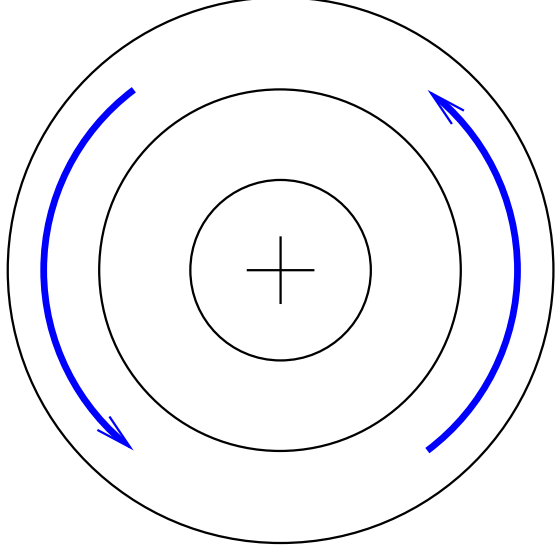
(Diamond and Kim, Phys Fl B 1991)



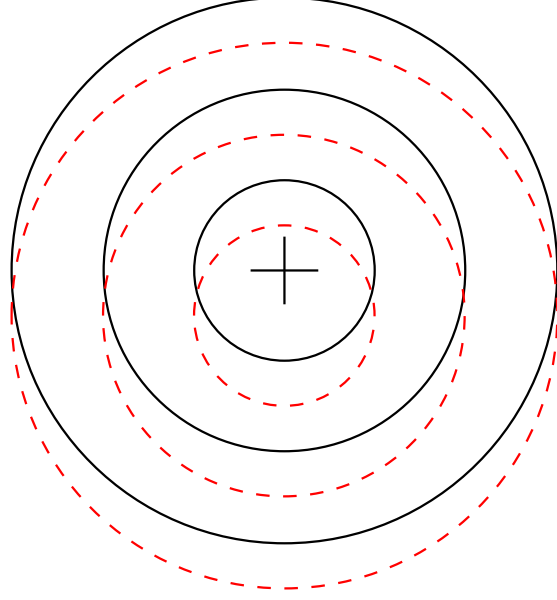
random alignment between zonal vorticity and Reynolds stress  
--> self generation/amplification of zonal flow layers

# Geodesic Acoustic Oscillation

(Winsor et al, Phys Fl 1968)



compression at top  
divergence at bottom



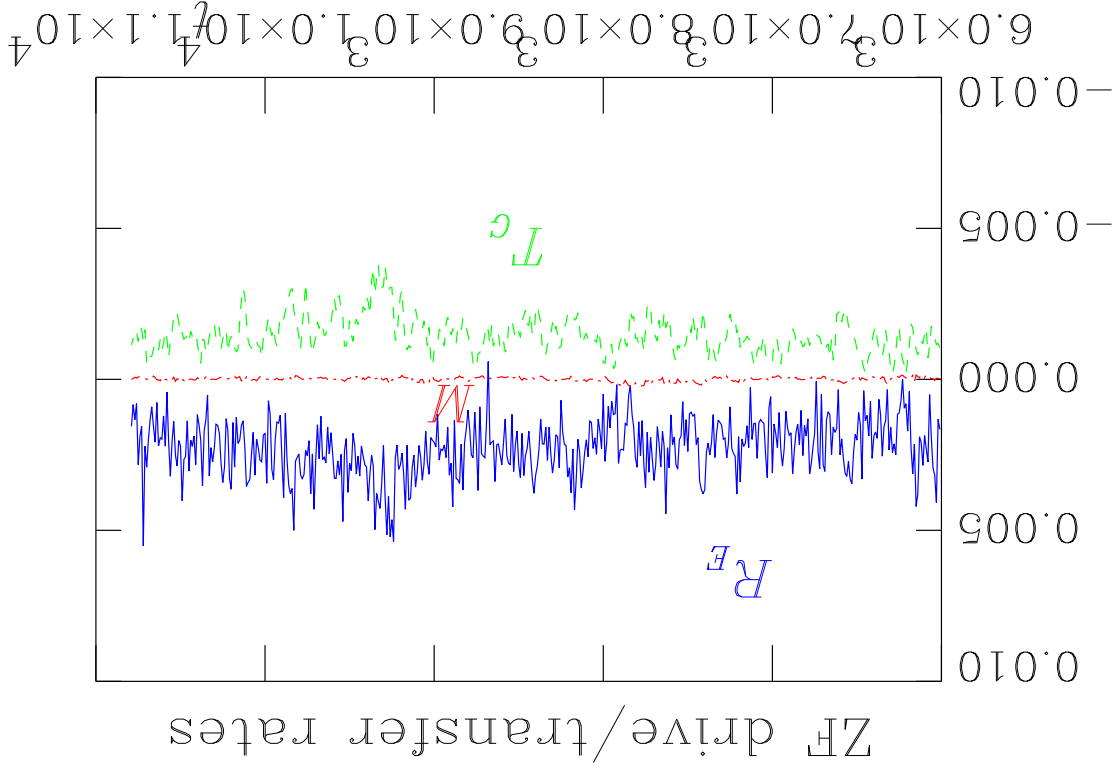
pressure sideband  
 $\langle p \sin \theta \rangle$

zonal flow exchanges conservatively with pressure sideband  
--> transfer pathway, equipartition



# Identification of Geodesic Transfer Effect

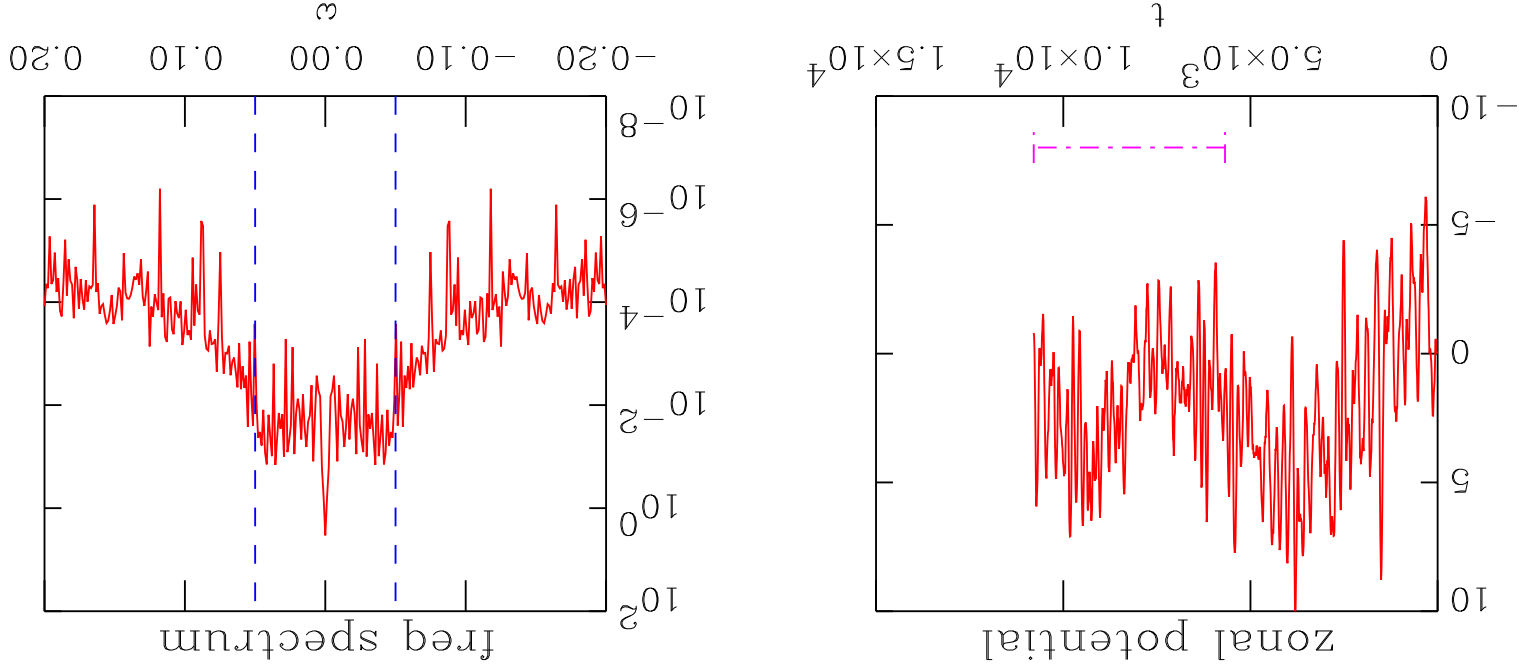
- measure Reynolds stress drive, geodesic transfer, and Maxwell stress in long term saturation



Reynolds stress / geodesic transfer in clear balance  
Maxwell stress is small

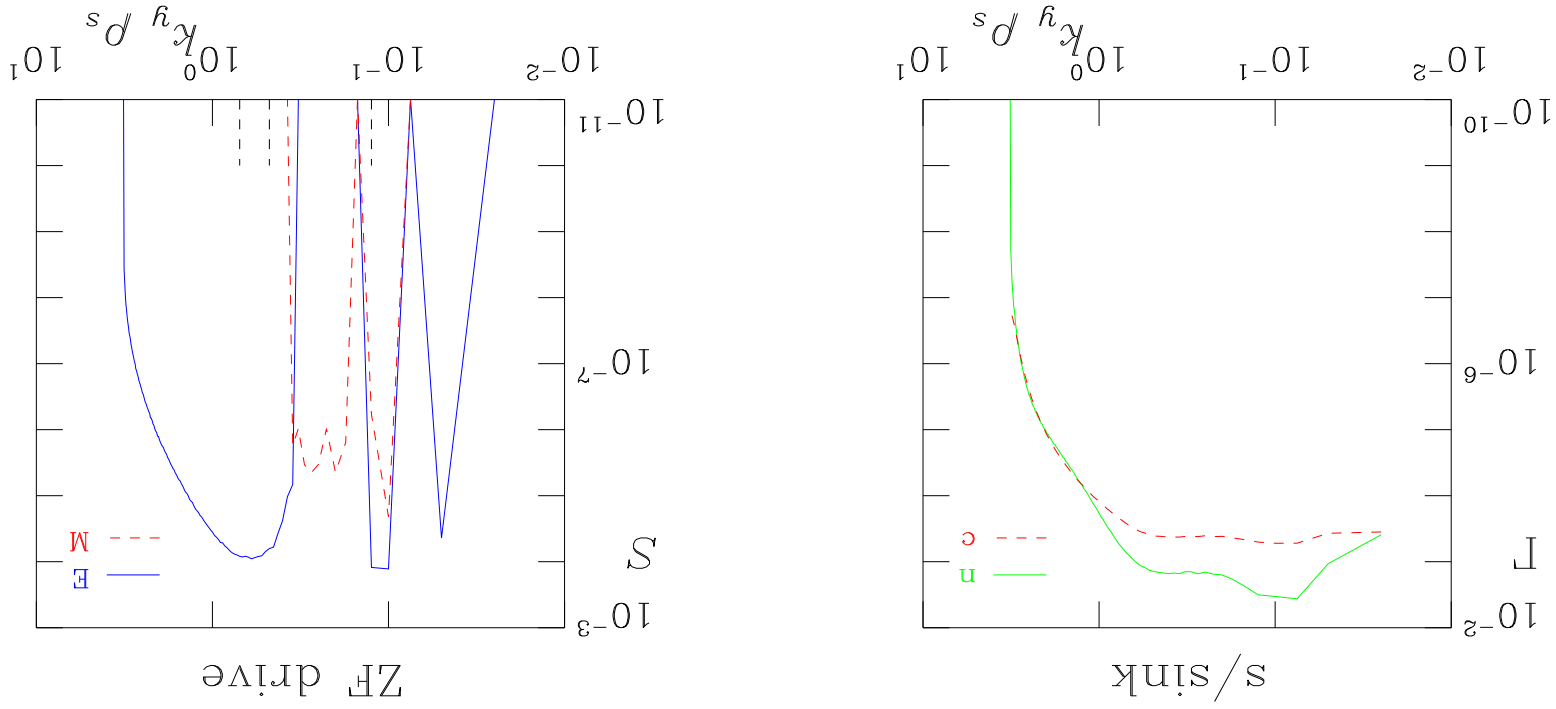
# Zonal Potential Frequency Spectrum

- zonal potential (check also flow, vorticity)
- geodesic acoustic oscillation small but visible



geodesic transfer, not oscillation, is main agent

# Transport and Reynolds Stress Spectra



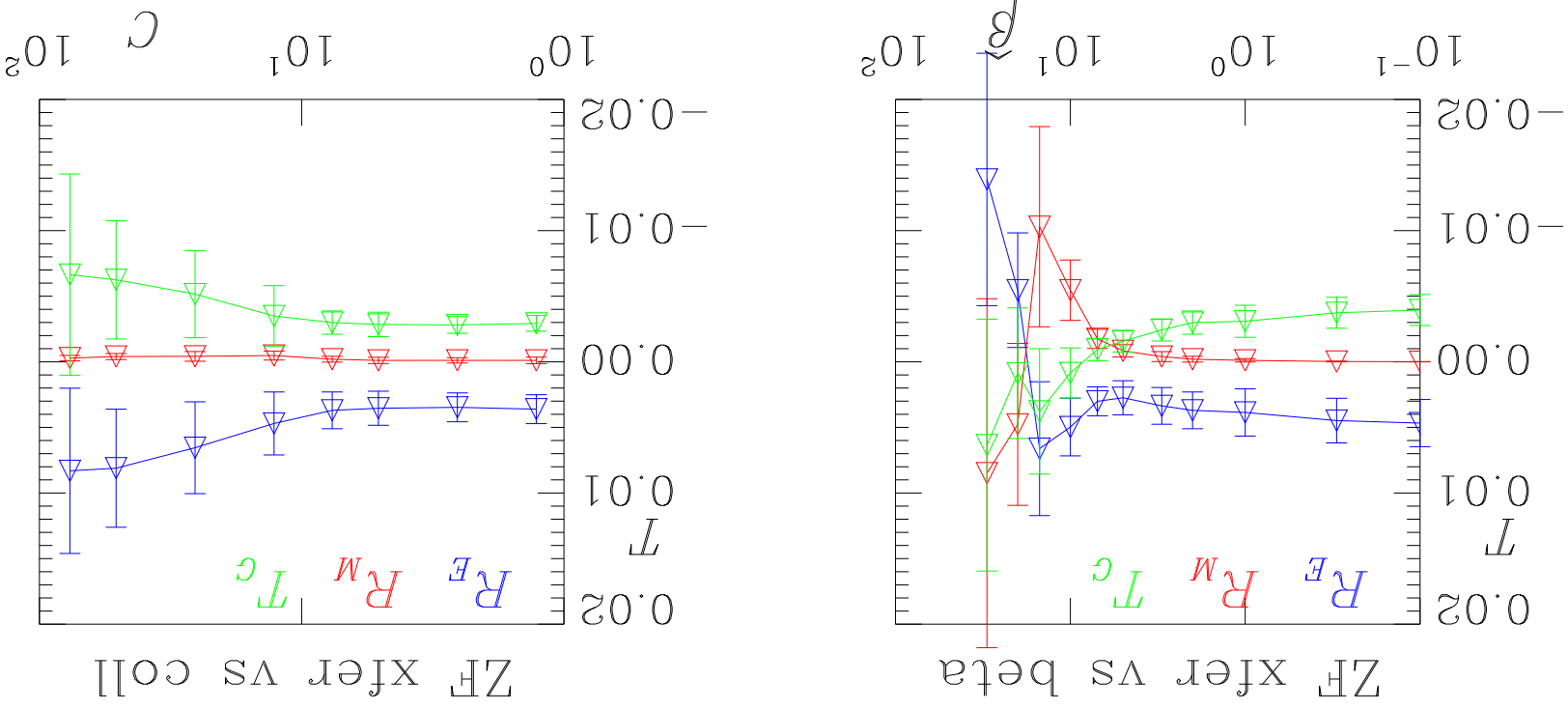
- transport: low (MHD mode) and high (DWT)  $k_y$  ranges
- Reynolds stress: 10, 20, and 50 percent lines

most of the transport and nearly all of the Reynolds stress done in the DWT spectral range near  $\rho_s$

# Scaling of Geodesic Coupling Effect

- basic case, ASDEX-Upgrade L-mode parameters ( $80 \text{ eV}$  and  $4.5 \times 10^{13} \text{ cm}^{-3}$ )

- sweep  $\beta$  and  $C$  around this case



geodesic coupling balances Reynolds stress in operational regime changes (PS-Alfvén oscillation) enter beyond ideal ballooning boundary

# Zonal Flow and Sideband Energetics (simplified)

- zonal flow energy ( $\Omega$  is vorticity)

$$\frac{\partial}{\partial t} \left\langle \frac{1}{2} \langle v_{\theta} \rangle^2 \right\rangle = \left\langle \Omega \right\rangle \left[ \langle v_x v_{\theta} \rangle - \frac{\beta}{1} \langle \rho_x \rho_{\theta} \rangle \right] - \omega_B \left\langle p_e \sin s \right\rangle \langle v_{\theta} \rangle$$

- the geodesic/Alfvén sideband system ( $\partial_x$  is radial turbulent flux,  $n_{\parallel}$  gives ion Landau losses)

$$\frac{\partial}{\partial t} \left\langle p_e \sin s \right\rangle^2 = 2 \left\langle \frac{\partial p_e}{\partial x} \sin s \right\rangle \langle \partial_x \sin s \rangle + \omega_B \left\langle p_e \sin s \right\rangle \langle v_{\theta} \rangle$$

$$- \omega_B \left\langle p_e \sin s \right\rangle \left\langle \frac{\partial p_e}{\partial x} \right\rangle - 2 \left\langle p_e \sin s \right\rangle \langle (J_{\parallel} - n_{\parallel}) \cos s \rangle$$

$$\frac{\partial}{\partial t} \left\langle \beta^{-1} \langle B_{\theta} \cos s \rangle^2 + \mu \langle J_{\parallel} \cos s \rangle^2 \right\rangle = 2 \left\langle (p_e - \phi) \sin s \right\rangle \langle J_{\parallel} \cos s \rangle - 2C \left\langle J_{\parallel} \cos s \right\rangle^2$$

$$\frac{\partial}{\partial t} \left\langle v_{\theta} \sin s \right\rangle^2 = 2 \left\langle \Omega \sin s \right\rangle \langle v_x v_{\theta} \sin s \rangle + \omega_B \left\langle \phi \sin s \right\rangle \left\langle \frac{\partial p_e}{\partial x} \right\rangle + 2 \left\langle \phi \sin s \right\rangle \langle J_{\parallel} \cos s \rangle$$

- the thermal reservoir (background pressure)

$$\frac{\partial}{\partial t} \left\langle \frac{1}{2} \langle p_e \rangle^2 \right\rangle = \left\langle \partial_x \partial \right\rangle \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \omega_B \left\langle \phi - p_e \right\rangle \left\langle \frac{\partial p_e}{\partial x} \right\rangle$$

## Zonal Flow and Sideband Dynamics (simplified)

- zonal potential  $\langle \phi \rangle$ , flow  $\langle v_y \rangle = \partial \langle \phi \rangle / \partial x$ , and vorticity  $\langle \Omega \rangle = \partial \langle v_y \rangle / \partial x$

- zonal vorticity (with nonlinear Reynolds and Maxwell stresses, geodesic curvature)

$$\frac{\partial}{\partial t} \langle \Omega \rangle = - \frac{\partial^2 \langle \phi \rangle}{\partial x^2} \left[ \langle v_x v_y \rangle - \frac{1}{2} \frac{\partial}{\partial x} \langle v_x^2 \rangle \right] - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

- the geodesic/Alfvén sideband system ( $\mathcal{Q}_x$  is radial turbulent flux,  $u_{||}$  gives ion Landau losses)

$$\frac{\partial}{\partial t} \langle p_e \sin s \rangle = \frac{\partial}{\partial x} \langle \mathcal{Q}_x \sin s \rangle + \frac{\omega_B}{2} \langle v_y \rangle - \frac{\omega_B}{2} \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \langle (J_{||} - u_{||}) \cos s \rangle$$

$$\frac{\partial}{\partial t} \langle \beta A_{||} + \mu J_{||} \rangle = \langle p_e \sin s \rangle - \langle \phi \sin s \rangle - C \langle J_{||} \cos s \rangle$$

$$\frac{\partial}{\partial t} \langle \Omega \sin s \rangle = - \frac{\partial^2 \langle \phi \rangle}{\partial x^2} \left( \langle v_x v_y \rangle - \frac{1}{2} \frac{\partial}{\partial x} \langle v_x^2 \rangle \right) \sin s - \frac{\omega_B}{2} \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \langle J_{||} \cos s \rangle$$

PS Alfvén

- the thermal reservoir (background pressure)

$$\frac{\partial}{\partial t} \langle p_e \rangle = \langle \mathcal{Q}_x \rangle + \omega_B \frac{\partial}{\partial x} \langle \phi \sin s \rangle - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

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PS adiabatic

$$\frac{\partial}{\partial t} \langle \beta A_{||} + \mu J_{||} \rangle \cos s = \langle p_e \sin s \rangle - \langle \phi \sin s \rangle - C \langle J_{||} \cos s \rangle$$

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PS Alfvén

$$\frac{\partial}{\partial t} \langle p_e \rangle = \langle \mathcal{Q}_x \rangle + \omega_B \frac{\partial}{\partial x} \langle \phi \sin s \rangle - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

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geodesic acoustic

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PS adiabatic

$$\frac{\partial}{\partial t} \langle \beta A_{||} + \mu J_{||} \rangle \cos s = \langle p_e \sin s \rangle - \langle \phi \sin s \rangle - C \langle J_{||} \cos s \rangle$$

$$\frac{\partial}{\partial t} \langle \Omega \sin s \rangle = \langle \frac{\partial^2 \langle \phi \rangle}{\partial x^2} \rangle \left( \langle v_x v_y \rangle - \frac{\beta}{1} \langle v_x v_y \rangle \right) \sin s - \frac{\omega_B}{2} \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \langle J_{||} \cos s \rangle$$

PS Alfvén

$$\frac{\partial}{\partial t} \langle p_e \rangle = \langle \hat{Q}_x \rangle + \omega_B \frac{\partial}{\partial x} \langle \phi \sin s \rangle - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

- the thermal reservoir (background pressure)



# Zonal Flow and Sideband Dynamics (simplified)

- zonal potential  $\langle \phi \rangle$ , flow  $\langle v_y \rangle = \partial \langle \phi \rangle / \partial x$ , and vorticity  $\langle \Omega \rangle = \partial \langle v_y \rangle / \partial x$

- zonal vorticity (with nonlinear Reynolds and Maxwell stresses, geodesic curvature)

$$\frac{\partial}{\partial t} \langle \Omega \rangle = - \frac{\partial^2 \langle \phi \rangle}{\partial x^2} \left[ \langle v_x v_y \rangle - \frac{\partial}{\partial x} \langle v_x v_y \rangle \right] - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

**DW TURB** geodesic acoustic

- the geodesic/Alfvén sideband system ( $\hat{Q}_x$  is radial turbulent flux,  $u_{||}$  gives ion Landau losses)

$$\frac{\partial}{\partial t} \langle p_e \sin s \rangle = \frac{\partial}{\partial x} \langle \hat{Q}_x \sin s \rangle + \frac{\omega_B}{2} \langle v_y \rangle - \frac{\omega_B}{2} \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \langle J_{||} \rangle (J_{||} - u_{||}) \cos s$$

**PS adiabatic**

$$\frac{\partial}{\partial t} \langle \beta A_{||} + \mu J_{||} \rangle \cos s = \langle p_e \sin s \rangle - \langle \phi \sin s \rangle - C \langle J_{||} \cos s \rangle$$

$$\frac{\partial}{\partial t} \langle \Omega \sin s \rangle = \langle \frac{\partial^2 \langle \phi \rangle}{\partial x^2} \rangle \left( \langle v_x v_y \rangle - \frac{\partial}{\partial x} \langle v_x v_y \rangle \right) \sin s - \frac{\omega_B}{2} \left\langle \frac{\partial p_e}{\partial x} \right\rangle - \langle J_{||} \cos s \rangle$$

**turb sink**

**PS Alfvén**

- the thermal reservoir (background pressure)

$$\frac{\partial}{\partial t} \langle p_e \rangle = \langle \hat{Q}_x \rangle + \omega_B \frac{\partial}{\partial x} \langle \phi \sin s \rangle - \omega_B \frac{\partial}{\partial x} \langle p_e \sin s \rangle$$

**sign of PS Alf xfer is reversed!**

# Summary and Outlook

- **Basic Results**
  - standard DALT turbulence
  - nonlinear self sustained drift wave instability
  - zonal flows and sidebands, global Alfvén transient → P-S current (2D eq)
  - zonal flows system reaches equilibrium on transport time scales
  - run to transport equilibrium very feasible for edge turbulence
- **Things GEM can do Now**
  - basic description of edge turbulence, **warm ions**
  - resolve  $p_z$ : note  $k_{\perp} p_z$  must go past unity to resolve DWT with  $T_e = T_i$
  - self consistent profiles with sources, turbulence
  - SOL/edge coupling, self consistent shear layer
  - experiment with sources, distribution thereof
- **Future Generalisation**
  - inhomogeneous equations (under development)
  - extension of gyrokinetic GENE code to edge turbulence
  - how much of this is needed to get pedestal (already know homogeneous models don't work)